

Si₃O₁₀ Groups in the Crystal Structure of Ardennite*

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Dedicated to Professor B. E. Warren for his 65th birthday

The rare mineral ardennite has the structural formula



Refined cell dimensions at 21 °C ($\lambda\text{Cu K}\alpha_1 = 1.54051 \text{ \AA}$) are 8.7126 ± 0.0008 , 5.8108 ± 0.0008 , $18.5214 \pm 0.0011 \text{ \AA}$; $Z = 2$, $Pnmm$, $D_{\text{calc}} = 3.74 \text{ g.cm}^{-3}$, $D_{\text{obs}} = 3.69\text{--}3.74 \text{ g.cm}^{-3}$. A total of 900 reflections were obtained on a Weissenberg-type automated Supper-Pace diffractometer supplemented by Weissenberg and precession film data. The structure was refined by least-squares method to a residual R of 7.0% including 110 unobserved reflections which were given a value of $F(\text{minimum})/\sqrt{3}$. The crystal structure can be derived from that of epidote by Ito-type twinning; it leads to the Si₃O₁₀ groups, which strengthen the bond chains along c . As in epidote, chains of octahedra sharing edges extend along b . They are here connected by the SiO₄ and Si₃O₁₀ groups. The cleavage is given as (001) in the literature for both epidote and ardennite; we determined it to be (100) for the two minerals. (As,V)O₄ tetrahedra do not share corners with SiO₄ tetrahedra. The polyhedra around Mn and especially (Mn,Ca) are irregular, suggesting abnormal magnetic behavior for ardennite. Bond lengths and temperature factors are normal. The average metal-oxygen distances are: for Mn 2.198 (5 nearest), for (Mn,Ca) 2.239 (5 nearest), for Al 1.911, for (Mg,Al,Fe) 2.011, for (As,V) 1.698, for Si 1.631 Å. The angle Si-O-Si is 126°.

Introduction

Ardennite is a rare arsenic-vanadium silicate, which has been found so far in the Ardennes, Belgium, in Piemont, Italy, and in the Jhabua district, India. The possibility that As and V might substitute for Si aroused our interest in the crystal structure. Professor J. D. H. Donnay collected a specimen at the type locality, Salm-Château, Ardennes, Belgium, which had been reopened on the occasion of the Liège meeting of the French Mineralogical Society in 1964. Other mineralogists did likewise, and as a result, Dr M. Semet and Professor J. Moreau (Semet & Moreau, 1965), of the University of Louvain, published a detailed account of ardennite mineralogy, with two new chemical analyses (Table 1) and a summary of previous work. They generously discontinued their own structure work when they heard of our efforts. A chemical analysis of a bulk sample (4.3 g) of our specimen was performed by C. O. Ingamells (Table 1).

Experimental procedure

The crystal chosen for data collection on the Weissenberg automated diffractometer has cell dimensions

$a = 8.7126 \pm 0.0008$, $b = 5.8108 \pm 0.0008$, $c = 18.5214 \pm 0.0011 \text{ \AA}$, at 21 °C. These were determined from back-reflection Weissenberg photographs about the a and b

Table 1. *Recent analyses of ardennite from Salm-Château, Belgium**

	1	2	3
SiO ₂	28.20	27.85	28.14
Al ₂ O ₃	23.80	23.55	23.22
MnO	26.75	25.70	25.33
Mn ₂ O ₃	—	—	0.00
As ₂ O ₅	11.25	13.25	9.85
V ₂ O ₅	1.86	0.89	0.82
Fe ₂ O ₃	1.22	0.86	1.50
FeO	0.00	0.00	0.00
MgO	2.30	2.30	3.83
CaO	1.17	1.17	1.50
CuO	0.57	0.26	0.46
Cr ₂ O ₃	—	—	0.12
Misc.†	—	—	0.11
F ₂	—	—	0.14
H ₂ O ⁺	5.10	5.10	5.04
H ₂ O ⁻	0.02	0.02	0.00
Total	101.26	100.95	100.06
Less O = F			0.06
			100.00

* 1, 2. Chemical analyses by O. Vandenberg, in Semet & Moreau (1965). 3. Chemical analysis by C. O. Ingamells (present paper).

† TiO₂, 0.00; SnO₂, 0.03; CoO, 0.01; NiO, 0.02; ZnO, 0.04; SrO, 0.00; BaO, 0.00; Na₂O, 0.01; K₂O, 0.00 (Sum = 0.11).

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axes, with Cu $K\alpha$ radiation ($\lambda_{\alpha_1}=1.54051$, $\lambda_{\alpha_2}=1.54434$ Å), the measurements of which were treated by the Burnham LCLSQ least-squares refinement program. The crystal size is $0.05 \times 0.08 \times 0.10$ mm. The density, calculated for 2 formula units per cell with the help of Ingamells's analysis (Table 1) to obtain the formula weight, is 3.74 g.cm^{-3} . Densities measured on the Berman balance range from 3.69 to 3.74 g.cm^{-3} for the crystals used in the electron-probe analysis. Semet & Moreau (1965, p. 559) concluded that 3.747 g.cm^{-3} should be considered the best experimental density for their ardennite sample. Systematic X-ray extinctions gave the aspect Pn^{**} . Statistical analyses of F and E values indicate the presence of a center of symmetry leading to the space group $Pnmm$ (D_{2h}^{13}); for example $\langle |E| \rangle$ equals 0.788 and $\langle |E^2 - 1| \rangle$ equals 1.061 as compared with the theoretical values 0.798 and 0.968 , respectively, for centrosymmetric space groups (Karle & Karle, 1963) and 0.89 and 0.74 , respectively, for non-centrosymmetric space groups (Karle & Karle, 1966a).

Our choice of axes is different from that of Semet & Moreau (1965); it is meant to stress the structural relation between epidote and ardennite, first pointed out by Strunz (1935), which will be discussed below. In our setting, the unique, easy, and perfect cleavage is (100), which is to be expected from the structure as will be shown below. Since this finding disagrees with the data in the literature, where the cleavage is reported to be perpendicular to the 18.52 Å translation, we checked it repeatedly. Cleavage fragments were mounted with the cleavage face in polar position on the two-circle goniometer. Precession photographs always showed the $E-W$ line to correspond to the 8.71 Å translation.

The intensities were measured on the diffractometer, in Cu $K\alpha$ radiation, from five b -axis levels containing 898 independent reflections, 275 of which were too weak to be observed. Weissenberg photographs, taken with Cu $K\alpha$ radiation and 12-hour exposures, permitted visual estimation of an additional 165 weak reflections, leaving only 110 unobserved ones. Two more reflections, 020 and 040, were obtained by visual estimation from a (001)₀ precession film taken with Mo $K\alpha$ radiation, thus bringing the total number of reflections to 900 (Table 5). The intensities were corrected for Lorentz and polarization factors; Burnham's GNABS program was used to eliminate absorption errors. An estimate of the accuracy of the observed F values, based on the scatter of 13 sets of symmetry-related strong and medium-strong test reflections, indicates a standard deviation for a single intensity measurement ranging from 3.4 to 25.0%, with an average value of 9.4%, so that an R value of 4.7% is the best we should hope for. A striking intensity feature is the relative faintness of all reflections with $h+k=2n+1$, indicating a pseudo C -centering of the structure. Reflections $10l$; $14l$; $0,2n+1,2n+1$ are especially weak. This pseudo C -centering is also evident on a three-dimensional Patterson synthesis, where the peak at $\frac{1}{2}\frac{1}{2}0$ reaches half the height of the origin peak.

Structure determination

A comparison of the ardennite and epidote cell dimensions is indicative of an Ito-type twin relation between the two structures: a is equal to 8.71 Å for ardennite and 8.88 for epidote; b equals 5.81 and 5.63 , respectively; c equals 18.52 for ardennite, is nearly equal to the $(2c \sin \beta)$ value of 18.32 for epidote (calculated from $c=10.22$ and $\beta=116^\circ 20'$, values we obtained for epidote from the Strathcona Mine, Sudbury area, Ontario, Canada).

One of us (Donnay, 1965, 1966) tried long and unsuccessfully to obtain a reasonable trial structure, beginning with the known epidote structure determined by Belov & Rumanova (1953) and Ito, Morimoto & Sadanaga (1954). The Patterson synthesis was calculated for us by Dr W. Fischer at the University Computing Center in Marburg, Germany. When analyzed in a manner analogous to the one used by Ito *et al.* (1954), this Patterson synthesis showed all the cations except (Mg, Al) but did not bring out the oxygen framework of tetrahedra and octahedra.*

We then decided to try the Karle-Karle symbolic addition procedure (Karle & Karle, 1963, 1966b), which had not yet been used on any inorganic compound containing heavy atoms. We obtained advice and generous computing aid from Dr I. Karle, who used the programs written in her laboratory to obtain a calculation of the E values and a listing of all possible combinations E_h, E_k, E_{h+k} with $E > 1.5$. A great number of strong relations between $E_{h2l}, E_{\bar{h}2l}, E_{040}$, and $E_{hk6}, E_{\bar{h}\bar{k}6}, E_{0012}$ proved the signs of 040 and 0012 to be positive. Three more signs were chosen to fix the origin and five more unknown signs were included in the unit factors a to e (Table 2). After two cycles of hand calculations with the Σ_2 equation, 30 signs were known; c was found to be negative and e positive. The third cycle gave 15 more signs, confirmed a to be

Table 2. Starting set of reflections used for statistical method

h	k	l	$ E $	Sign	
4	2	5	3.93	+	} from choice of origin
7	2	5	2.66	+	
1	1	5	2.54	+	
0	4	0	3.98	+	} from strong relations
0	0	12	2.84	+	
8	0	2	2.63 <i>a</i>		a becomes +1
4	2	6	3.80 <i>b</i>		b doubtful
4	2	7	3.54 <i>c</i>		c becomes -1
2	2	5	2.94 <i>d</i>		d becomes +1
5	3	10	2.82 <i>e</i>		e becomes +1

* After the structure was solved, the coordinates of the cations as found in the Patterson synthesis were used for structure-factor calculations. In this calculation, 335 signs out of 369 for the reflections with $F_o > 36.0$ were found to be correct. At the time when we had only the data from the Patterson synthesis, however, we did not trust our results because we could not place the expected anion polyhedra around the observed cation sites.

Table 3. *Fractional atomic coordinates in ardennite structure*

Values obtained from Patterson function, P ; from E map, followed by q_E maximum; final refined coordinates with standard deviation and isotropic temperature factor B from last isotropic refinement.

Atom and position occupied	From P	From E	q_E Max.	Final	B (\AA^2)
(Mn,Ca) 4(f)	0.950 $\frac{1}{4}$ 0.154	0.950 $\frac{1}{4}$ 0.155	396	0.9441 \pm 2 $\frac{1}{4}$ 0.1554 \pm 1	0.95 \pm 5
Mn 4(f)	0.400 $\frac{3}{4}$ 0.152	0.400 $\frac{3}{4}$ 0.160	392	0.3956 \pm 3 $\frac{3}{4}$ 0.1600 \pm 1	0.95 \pm 5
Al(1) 4(c)	0 0 0	0 0 0	214	0 0 0	0.62 \pm 7
Al(2) 4(d)	$\frac{1}{2}$ 0 0	$\frac{1}{2}$ 0 0	252	$\frac{1}{2}$ 0 0	0.76 \pm 7
(Mg,Al,Fe) 4(e)		0.675 -0.002 $\frac{1}{4}$	213	0.6787 \pm 4 -0.0016 \pm 8 $\frac{1}{4}$	0.96 \pm 7
Si(1) 4(f)	0.750 $\frac{3}{4}$ 0.096	0.767 $\frac{3}{4}$ 0.100	340	0.7641 \pm 4 $\frac{3}{4}$ 0.0970 \pm 2	0.37 \pm 6
Si(2) 4(f)	0.267 $\frac{1}{4}$ 0.091	0.270 $\frac{1}{4}$ 0.097	264	0.2766 \pm 4 $\frac{1}{4}$ 0.0951 \pm 2	0.59 \pm 7
Si(3) 2(a)	0.579 $\frac{1}{4}$ $\frac{1}{4}$	0.333 $\frac{1}{4}$ $\frac{1}{4}$	205	0.3412 \pm 5 $\frac{1}{4}$ $\frac{1}{4}$	0.66 \pm 9
(As,V) 2(b)	0.046 $\frac{3}{4}$ $\frac{1}{4}$	0.042 $\frac{3}{4}$ $\frac{1}{4}$	633	0.0426 \pm 2 $\frac{3}{4}$ $\frac{1}{4}$	0.45 \pm 5
O(1) 8(g)		0.117 0.025 -0.078	104	0.1284 \pm 6 0.0265 \pm 13 -0.0828 \pm 3	0.82 \pm 11
O(2) 8(g)		0.622 -0.008 -0.083	130	0.6241 \pm 6 -0.0207 \pm 13 -0.0833 \pm 3	0.81 \pm 11
O(3) 4(f)		0.133 $\frac{1}{4}$ 0.042	84	0.1127 \pm 9 $\frac{1}{4}$ 0.0526 \pm 4	0.86 \pm 16
O(4) 4(f)		0.617 $\frac{3}{4}$ 0.040	123	0.6169 \pm 10 $\frac{3}{4}$ 0.0415 \pm 4	0.91 \pm 16
O(5) 4(e)		0.950 -0.050 $\frac{1}{4}$	114	0.9299 \pm 10 -0.0068 \pm 21 $\frac{1}{4}$	1.02 \pm 15
O(6) 4(e)		0.438 0.042 $\frac{1}{4}$	88	0.4431 \pm 9 0.0146 \pm 20 $\frac{1}{4}$	1.06 \pm 16
O(7) 4(f)		0.675 $\frac{3}{4}$ 0.175	105	0.6817 \pm 10 $\frac{3}{4}$ 0.1776 \pm 5	1.19 \pm 16
O(8) 4(f)		0.225 $\frac{1}{4}$ 0.182	81	0.2232 \pm 9 $\frac{1}{4}$ 0.1805 \pm 4	0.78 \pm 16
O(9) 4(f)		0.200 $\frac{3}{4}$ 0.182	58	0.1540 \pm 11 $\frac{3}{4}$ 0.1763 \pm 5	1.36 \pm 17

Table 3 (cont.)

Atom and position occupied	From <i>P</i>	From <i>E</i>	<i>q_E</i> Max.	Final	<i>B</i> (Å ²)
OH(1)		0.120		0.1128 ± 10	
4(<i>f</i>)		$\frac{3}{4}$	101	$\frac{3}{4}$	0.95 ± 16
		0.028		0.0323 ± 5	
OH(2)		0.622		0.6149 ± 9	
4(<i>f</i>)		$\frac{1}{4}$	162	$\frac{1}{4}$	0.35 ± 14
		0.038		0.0356 ± 4	
OH(3)		0.700		0.6973 ± 10	
4(<i>f</i>)		$\frac{1}{4}$	132	$\frac{1}{4}$	0.91 ± 17
		0.188		0.1817 ± 4	

Table 4. Anisotropic temperature factors for ardennite

The β_{ij} of the expression

$$\exp [-(\beta_{11}h^2 + \beta_{22}k^2 + \beta_{33}l^2 + \beta_{23}2kl + \beta_{31}2lh + \beta_{12}2hk)]$$

are converted into B_{ij} in Å², which are directly comparable with the B values of Table 3.

	B_{11}	B_{22}	B_{33}	B_{12}	B_{13}	B_{23}
(Mn,Ca)	1.18 ± 8	0.89 ± 15	0.82 ± 7	0	0.16 ± 6	0
Mn	1.03 ± 7	1.23 ± 15	0.79 ± 7	0	-0.15 ± 6	0
Al(1)	0.64 ± 11	0.78 ± 20	0.52 ± 11	-0.10 ± 14	0.04 ± 9	-0.05 ± 13
Al(2)	0.82 ± 12	0.91 ± 21	0.61 ± 11	0.06 ± 15	-0.02 ± 10	-0.01 ± 13
(Mg,Al,Fe)	1.29 ± 13	0.95 ± 22	0.76 ± 12	-0.08 ± 15	0	0
Si(1)	0.38 ± 11	0.56 ± 23	0.33 ± 11	0	-0.02 ± 9	0
Si(2)	0.56 ± 12	0.94 ± 24	0.47 ± 12	0	-0.03 ± 9	0
Si(3)	0.75 ± 16	0.80 ± 33	0.50 ± 16	0	0	0
(As,V)	0.63 ± 7	0.27 ± 14*	0.45 ± 6	0	0	0
O(1)	1.30 ± 21	0.88 ± 42	0.48 ± 21	-0.24 ± 23	-0.17 ± 17	-0.07 ± 21
O(2)	0.67 ± 20	1.03 ± 40	0.75 ± 21	0.20 ± 23	0.05 ± 17	0.28 ± 13
O(3)	0.72 ± 31	1.92 ± 57	0.48 ± 31	0	-0.14 ± 25	0
O(4)	0.96 ± 31	1.69 ± 57	0.50 ± 30	0	-0.01 ± 25	0
O(5)	1.39 ± 31	1.00 ± 50	0.88 ± 30	-0.13 ± 36	0	0
O(6)	0.91 ± 30	1.43 ± 53	1.22 ± 30	-0.03 ± 35	0	0
O(7)	1.01 ± 32	1.71 ± 63	0.74 ± 32	0	0.5 ± 27	0
O(8)	0.91 ± 30	0.33 ± 57	1.01 ± 32	0	-0.26 ± 25	0
O(9)	1.79 ± 36	1.26 ± 62	1.00 ± 33	0	0.35 ± 30	0
OH(1)	0.90 ± 32	1.08 ± 56	0.91 ± 33	0	0.24 ± 26	0
OH(2)	0.67 ± 26	0.27 ± 53*	0.27 ± 29*	0	-0.13 ± 22	0
OH(3)	0.74 ± 31	0.91 ± 60	1.06 ± 34	0	0.42 ± 25	0

* Values were kept constant, because they tended to become negative.

positive, and indicated d to be very probably positive. There was no hint as yet of the correct sign of b . After the fifth cycle 95 out of 118 signs with $E > 1.5$ were known, still leaving b undetermined. Then, for both possible signs of b , three more Σ_2 cycles for all 168 E values > 1.3 were calculated by computer, immediately followed by a three-dimensional E map. In the case of $b = +1$, 159 signs could be determined, and q_{\min} in the E map was -178 ; whereas for $b = -1$, 150 signs could be determined, and q_{\min} was -185 . The latter E map made no sense at all, but the one calculated with $b = +1$ showed all the atoms, the cations being nearly in the same places as in the Patterson synthesis. The peak heights gave a good indication of the chemical nature of the atoms (Table 3). By checking against the final calculated signs, only one out of these 159 signs was found to be wrong (that of 702).

Refinement of coordinates was carried out with the modified Busing-Martin-Levy least-squares program at The Johns Hopkins University Computing Center. For atomic scattering factors we used As³⁺ from *International Tables for X-ray Crystallography*, Vol. III, p. 206; Mn⁺, Ca⁺, Fe⁺, Al⁺, Mg²⁺, V³⁺, Si³⁺, and O⁻ calculated by Hanson, Herman, Lea & Skillman (1966). With isotropic temperature factors and F (unobserved) set equal to $(1/\sqrt{3})F$ (minimum), the residual $R = \sum |F_o| - |F_c| / \sum |F_o|$ dropped from 18.0 to 8.4%; continuing with anisotropic temperature factors the residual dropped to 7.9%. We experimented with refinement of the multiplier for various scattering factors but obtained no significant improvement.

Next we introduced anomalous dispersion corrections for all atoms, using the values of Dauben & Templeton (1955) and the SFLS5 least-squares pro-

gram of Prewitt. The consequent deviations from 0 (+) and 500 (-) for the phase angle, expressed in millicycles, run as high as 200 for several reflections (Table 5). Unobserved reflections are given the value $|F_{\min}|/3$. The R value with isotropic temperature factors (Table 3) for all 900 reflections dropped to 7.5% and, continuing with anisotropic temperature factors (Table 4) came to 7.0%. For the 240 strongest reflections with $F_o \geq 50$ an R value of 4.7% is obtained, in good agreement with the accuracy of our intensity measurements.

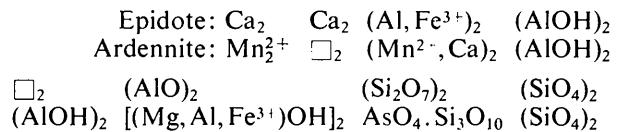
A comparison of atomic coordinates given by the Patterson map, the E map, and the final refinement (Table 3) shows the close agreement between the initial coordinates obtained from the E map and the final, refined ones. The bond lengths and their Pauling bond strengths (Table 6), as well as oxygen-oxygen approaches (Table 7), all fall within normal limits. Pauling bond strengths have been related roughly to bond lengths for all the cations except aluminum and silicon.

Discussion

The new and unexpected feature of the ardennite structure consists in the Si_3O_{10} groups of three tetrahedra, all pointing the same way and not alternating as in infinite chains, possessing symmetry $2mm$, and Si-O-Si bond angles of 126° . The usual value for this angle is close to 135° in infinite chains. This tritetrahedral group occurs jointly with individual SiO_4 groups in the ratio 1:2. Liebau (1962) surveyed the known silicate structures and stated the following rule: *tetrahedral groups that differ by more than one in the number of shared corners should not occur jointly in a silicate structure*. Professor F. Liebau (personal communication)

pointed out that ardennite is the third known exception to this rule; the other two are prehnite and zunyite.

The structural relationship with epidote is close indeed (Fig. 1): a mirror (001), passing through the origin of the cell, reflects the projected epidote structure in such a way that only very minor rearrangements in the immediate neighborhood of the mirror are needed to give ardennite (Fig. 2). In all the trial structures that we considered, we overlooked the possibility of placing one of the tetrahedra of the Si_2O_7 group on the 'Ito mirror', which thus reflects the second tetrahedron of the Si_2O_7 group on the other side, producing the Si_3O_{10} group. The formulae for the cell content of epidote and half the cell content of ardennite can be written so as to show their structural relation:



where \square is a vacant site.

The epidote structure is composed of five-membered 'mixed rings' of the 'chair form' consisting of an aluminum octahedron, a silicon tetrahedron, another aluminum octahedron, and two linked silicon tetrahedra. In ardennite, two of these mixed rings, in addition to sharing their octahedra, share the central tetrahedron of the Si_3O_{10} group, forming a large bow-tie of two five-membered rings. Inversion through the centers of symmetry at the Al sites in both structures reproduces the individual rings or bow-ties and extends the chains along c . Combined with the chains of octahedra along b , the structures can be considered to consist of thick sheets parallel to (100). Neither mineral species can be

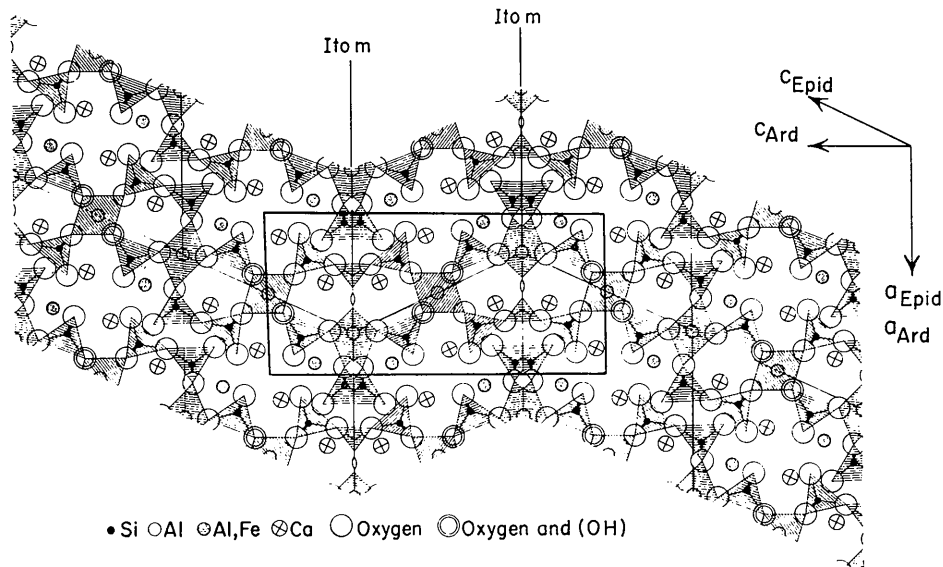


Fig. 1. The epidote structure (ten cells in all, after Ito *et al.*, 1954), projected on (010). The Ito mirrors, through the cell origin, are a distance $d(001)$ apart. The outline of the orthorhombic cell related to that of ardennite, also projected on (010), is shown with heavy lines.

Table 6. Bond lengths (Å), each with its standard deviation and estimated Pauling bond strength

Anions	Cations									ΣII‡
	(Mn,Ca)	Mn	Al(1)	Al(2)	(Mg,Al,Fe)	Si(1)	Si(2)	Si(3)	(As,V)	
O(1)	2.189 ± 7 0.5*		1.904 ± 7 0.5*			1.623 ± 8 1.0*				2.0
O(2)		2.127 ± 7 0.5*		1.888 ± 7 0.5*			1.604 ± 8 1.0*			2.0
O(3)	(2.403 ± 8) 0.0		2.006 ± 8 0.5†				1.631 ± 9 1.0			2.0
O(4)				1.934 ± 9 0.5†		1.644 ± 9 1.0				2.0
O(5)	2.305 ± 10 0.3†				2.189 ± 10 0.3				1.721 ± 10 1.1*	2.0
O(6)		2.305 ± 10 0.25†			2.055 ± 10 0.4			1.631 ± 10 1.0*		1.9
O(7)		(2.514 ± 9) 0.0			1.970 ± 10 0.45†	1.657 ± 10 1.0				1.9
O(8)	(2.473 ± 8) 0.0						1.649 ± 9 1.0	1.647 ± 9 1.0*		2.0
O(9)		2.127 ± 9 0.5							1.675 ± 9 1.4*	1.9
OH(1)			1.853 ± 9 0.5†							1.0
OH(2)				1.883 ± 8 0.5†						1.0
OH(3)	2.207 ± 8 0.4				1.940 ± 9 0.45†					1.3
ΣI‡	2.0	2.0	3.0	3.0	2.5	4.0	4.0	4.0	5.0	
Cation coordination number	5	5	6	6	6	4	4	4	4	
Average bond length	2.239 (2.296)	2.198 (2.250)	1.921	1.902	2.011	1.637	1.622	1.639	1.698	

* Two bonds for cation, one for anion.

† Two bonds for cation, two for anion.

‡ ΣI = sum of Pauling bonds emanating from cation. ΣII = sum of Pauling bonds reaching anion.

Table 7. Oxygen approaches less than 3 Å in ardennite

Standard deviations = 0.01 Å.

As in Fig. 2, the primes represent atoms in sites that are transformed as follows: $xyz' = \bar{x}, \bar{y}, \bar{z}$; $xyz'' = 1 - x, \bar{y}, \bar{z}$; $xyz''' = x, y, \frac{1}{2} - z$; $xyz^{iv} = 1 - x, 1 - y, \bar{z}$; $xyz^v = 1 - x, \frac{1}{2} + y, \bar{z}$; $xyz^{vi} = 1 - x, \frac{1}{2} + y, \bar{z}$; $xyz^{vii} = x, \frac{1}{2} - y, z$; $xyz^{viii} = x - 1, 1 + y, z$; $xyz^{ix} = x - 1, \frac{1}{2} - y, z$.

Around Al(1)		Around Al(2)		Around (Mg,Al,Fe)		Empty octahedron around ($\frac{1}{4}0\frac{1}{2}$)	
Two O(1)–O(3)	2.83	Two O(2)–O(4)	2.67	Two O(5)–O(7)	2.91	3 sides as in oct. Al(1)	
Two O(1)–O(3')	2.70	Two O(2)–O(4'')	2.73	Two O(5)–OH(3)	2.82	3 sides as in oct. Al(2)	
Two O(1)–OH(1)	2.67	Two O(2)–OH(2)	2.71	Two O(6)–O(7)	2.91	1 side as in tet. Si(1)	
Two O(1)–OH(1')	2.64	Two O(2)–OH(2'')	2.63	Two O(6)–OH(3)	2.89	1 side as in tet. Si(2)	
Two O(3)–OH(1)	2.93	Two O(4)–OH(2)	2.91	Two O(7)–OH(3)	2.91	O(1)–OH(2'')	2.89
Two* O(3)–OH(1')	2.52	Two* O(4)–OH(2'')	2.47	* O(7)–O(7''')	2.68	O(2'')–OH(1)	2.94
				* OH(3)–OH(3''')	2.53	O(3)–O(4'')	2.93
Average of 10	2.75	Average of 10	2.73	Average of 10	2.89	OH(1)–OH(2'')	2.69
Average of 12	2.71	Average of 12	2.69	Average of 12	2.84	Average of 4	2.86
Around Si(1)		Around Si(2)		Around Si(3)		Around (As,V)	
Two O(1 ^{iv})–O(1 ^v)	2.60	Two O(2'')–O(2 ^{vi})	2.66	O(6)–O(6 ^{vii})	2.74	O(5 ^{viii})–O(5 ^{ix})	2.83
Two O(1 ^{iv})–O(4)	2.68	Two O(2'')–O(3)	2.71	Four O(6)–O(8)	2.68	Four O(5 ^{viii})–O(9)	2.77
Two O(1 ^{iv})–O(7)	2.74	Two O(2'')–O(8)	2.61	O(8)–O(8''')	2.57	O(9)–O(9)	2.73
O(4)–O(7)	2.58	O(3)–O(8)	2.56	Average of 6	2.67	Average of 6	2.77
Average of 6	2.67	Average of 6	2.64				
Other distances:		O(1)–OH(3'')	2.87	O(9)–OH(1)	2.69		
		O(8)–O(9)	2.97	OH(2)–OH(3)	2.80		

* Shared edge.

expected to show a (001) cleavage that would cut through the strong bonds of the sheets. A perfect (001) cleavage is, however, reported for epidote throughout the literature and was discussed at length in the papers describing the epidote-structure determination. We tested the cleavage of epidote crystals from many localities and not once did we observe a (001) cleavage plane. If it exists at all, and is not just a misprint in the literature, it must be extremely difficult to produce. The (100) cleavage parallel to the chains, imperfect in epidote, is perfect in ardenite.

All the coordination polyhedra are normal except those about Mn^{2+} and (Mn^{2+}, Ca) for which a regular octahedral environment is expected. One of the six Mn-O bonds is 0.3 Å longer than the average bond length of the other five (Table 6), leading to a distorted octahedron around Mn. The five nearest bonds emanating from (Mn,Ca) have a configuration similar to that of the shortest five Mn-O bonds. There are, however, two next-nearest oxygen atoms only 0.2 Å farther away than the average (Mn,Ca)-(O,OH) bond length. Including them would give sevenfold coordination (six O and one OH). We may therefore expect ardenite to show unusual magnetic properties, and a study of its magnetic susceptibility down to liquid-helium temperature is under way.*

The original question, whether all the As and V ions that are present in ardenite substitute for Si, received

* Note added in proof.: - Dr F. Senftle and coworkers have observed antiferromagnetism below liquid nitrogen temperatures.

a negative answer. An $(As,V)O_4$ tetrahedron, sharing two of its corners with (Mg,Al,Fe) and (Mn,Ca) polyhedra and the other two with Mn polyhedra, exists in the structure. It is readily identified by its peak height in the E map (Table 3) and by the metal-oxygen bond lengths (Table 6). In addition, the assumption that all As and V are in the $2(b)$ space-group position, and the consequent choice of scattering factors, led to normal values for the temperature factors (Table 3). The slightly low B value of Si(1) suggests that this isolated Si(1) tetrahedron is the most likely one to incorporate any excess of As and V.

We therefore looked at the chemical data in the literature and transformed the analyses considered most reliable by Senet & Moreau (1965, p.560) to structural formulae for $(O+OH)=56$ (Table 8). As and V form a complete series of solid solutions. The amount of Ca substituting for Mn also varies within wide limits. Both (Mn,Ca) and (Mg,Al, etc.) positions show surprisingly large and unexplained variations from the expected values of 4 each, and their sum, instead of being 8, ranges from 6.90 to 8.65. In conclusion, it is fair to state that starting with the available chemical analyses one could hardly have arrived at the correct structural formula given above.

The temperature factors of the anions are normal (Table 3), except that of OH(2), which is much too low. Only complete substitution of F for OH(2) could possibly account for it, but the chemical analysis of our specimen (Tables 1 and 8) indicates that not more than 4 at.% of F is available for replacement of OH(2).

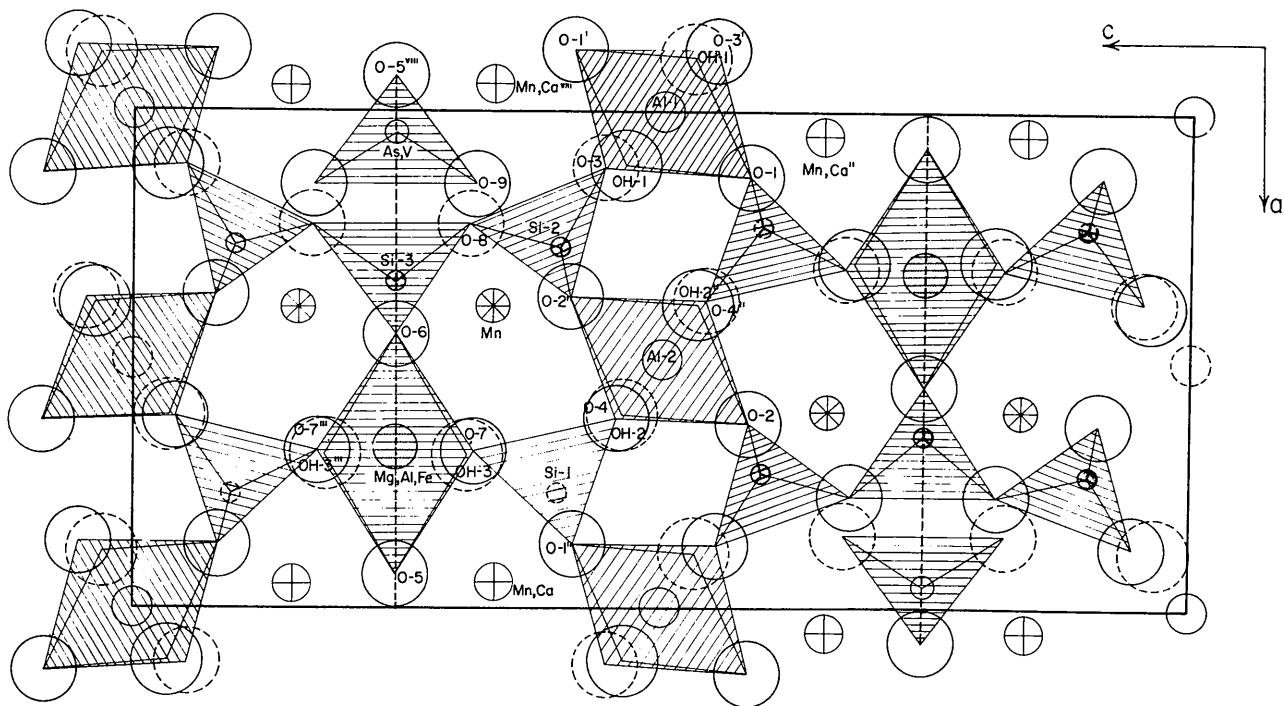


Fig. 2. Projection of refined ardenite structure on (010). Note that the coordinates given in Table 3 refer to an origin at the midpoint of the c repeat in this figure. Unprimed atoms refer to coordinates given in Table 3; primed atoms are defined in legend of Table 7.

Table 8. *Chemical analyses of ardennite written as structural formulae, so that (O+OH)=56*

All samples are from Salm-Château, Belgium, except No. 5, which comes from Ceres, Val d'Ala, Piemont, Italy. Only full analyses were used. Two decimal places are given throughout; they are not to be considered a reflection on the quality of the individual analysis. Each analysis contains 4 Mn, 8 Al, and 6 Si in addition to the amounts shown below.

	4(Mn, Ca)	4(Mg, Al, Fe ³⁺)	2(As, V)	4(Si, As)	12(OH)	O 44	Analyst	Reference
1	Mn 3.53 Ca 0.49 <u>4.02</u>	Mg 2.00 Al 0.60 Fe 0.40 Cu 0.12 <u>3.12</u>	As 1.81 V 0.19 <u>2.00</u>	Si 3.87	OH 11.79 F 0.16 <u>11.95</u>	O 44.05	C. O. Ingamells	Present paper
2	Mn 3.90 Ca 0.44 <u>4.34</u>	Mg 1.19 Al 1.77 Fe 0.32 Cu 0.15 <u>3.43</u>	As 1.80 V 0.20 <u>2.00</u>	Si 3.83 As 0.25 <u>4.08</u>	OH 11.83	O 44.17	O. Vandenberg	Semet & Moreau (1965)
3	Mn 3.50 Ca 0.43 <u>3.93</u>	Mg 1.18 Al 1.56 Fe 0.16 Cu 0.07 <u>2.97</u>	As 1.99 V 0.20 <u>2.19</u>	Si 3.60 As 0.40 <u>4.00</u>	OH 11.70	O 44.30	O. Vandenberg	Semet & Moreau (1965)
4	Mn 3.82 Ca 0.95 Fe ⁺² 0.27 <u>5.04</u>	Mg 1.04 Al 2.19 Fe ⁺³ 0.03 <u>3.26</u>	As 0.76 V 1.12 <u>1.88</u>	Si 4.00	OH 11.80	O 44.20	B. Gossner & H. Strunz	Gossner & Strunz (1932)
5	Mn 1.51 Ca 2.08 <u>3.59</u>	Mg 2.08 Al 1.15 Fe 0.81 Cu 0.20 <u>4.24</u>	As 0.87 V 0.06 P ⁺⁵ 0.04 Si 0.24 <u>1.21</u>	Si 4.00	OH 12.51	O 43.49	F. Zambonini	Zambonini (1922)
6	Mn 2.66 Ca 0.53 Fe ⁺² 0.19 <u>3.38</u>	Mg 2.68 Al 1.92 Fe ⁺³ 0.14 Cu 0.06 <u>4.80</u>	As 1.79 V 0.19 <u>1.98</u>	Si 3.62 Al 0.38 <u>4.00</u>	OH 12.08	O 43.92	W. Prandtl	Prandtl (1905)
7	Mn 3.87 Ca 0.81 <u>4.68</u>	Mg 1.56 Al 1.94 <u>3.50</u>	As 0.19 V 2.12 <u>2.31</u>	Si 3.69 As 0.31 <u>4.00</u>	OH 11.61	O 44.39	A. Bettendorff	von Lasaulx (1876)
8	Mn 4.28 Ca 0.59 <u>4.87</u>	Mg 0.74 Al 1.60 Fe 0.31 Cu 0.05 <u>2.70</u>	As 1.75 V 0.13 <u>1.88</u>	Si 3.85	OH 12.22	O 43.78	A. Bettendorff	von Lasaulx (1876)
9	Mn 3.43 Ca 1.09 <u>4.52</u>	Mg 2.07 Al 1.66 Fe 0.34 Cu 0.06 <u>4.13</u>	As 1.13 V 0.70 <u>1.83</u>	Si 3.69 Al 0.31 <u>4.00</u>	OH 11.81	O 44.19	F. Pisani	Pisani (1873)
10	Mn 3.67 Ca 0.68 <u>4.35</u>	Mg 1.76 Al 1.68 Fe 0.44 <u>3.88</u>	V 2.13	Si 4.35	OH 9.40	O 46.60	A. von Lasaulx & A. Bettendorff	von Lasaulx & Bettendorff (1873)
11	Mn 3.74 Ca 0.77 <u>4.51</u>	Mg 1.79 Al 1.74 Fe 0.51 <u>4.04</u>	V 2.12	Si 4.47	OH 9.46	O 46.54	A. von Lasaulx & A. Bettendorff	von Lasaulx & Bettendorff (1873)

The increased length of the *b* cell edge (5.81 Å) in ardennite over that found in other aluminum silicates with octahedral chains along *b* (5.45–5.63 in zoisite, 5.57 in kyanite, 5.61–5.64 in epidote, 5.66 in staurolite, 5.75 in allanite) is due to the substitution of the larger Mg and Fe ions for Al in about half of the octahedra that have centers on the mirror planes (001).

As can be seen on Fig. 2, a nearly regular but empty octahedron around the point $\frac{1}{4}00$ shares a face with each of two Al octahedra and an edge with each of two Si tetrahedra. The average length of the remaining four edges is 2.86 Å (Table 7), very comparable to the edge length of the other filled octahedra. The reason that no cation, M, is found at the center is ascribed to the short Al(1)–Al(2) separation, which would lead to M–Al distances of only 2.18 Å.

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The Configuration at C(13) in Labdanolic and Eperuic Acids by an X-ray Crystal Structure Analysis of the *p*-Bromophenacyl Ester of Labdanolic Acid

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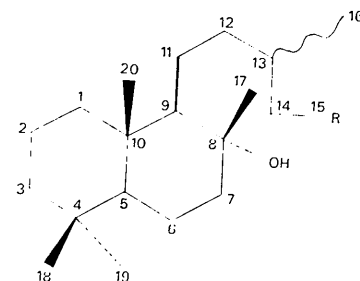
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The crystal structure of the *p*-bromophenacyl ester of labdanolic acid, determined by three-dimensional X-ray analysis, establishes the common configuration for labdanolic and eperuic acids at C(13) as (*S*) (Prelog's notation). The crystals are monoclinic, space group *P2*₁, with two molecules of C₂₈H₄₁O₄Br in a unit cell of dimensions *a* = 11.55, *b* = 6.05, *c* = 19.90 Å; β = 99.0°. The structure has been refined by least-squares methods to a final *R* of 11.4% for 1672 reflexions.

Introduction

Labdanolic acid, C₂₀H₃₆O₃, was first isolated by Cocker, Halsall & Bowers, 1956) from Spanish gum labdanum obtained from *Cistus ladaniferus* (rockrose). The structure of this new bicyclic diterpene hydroxy-acid was determined in a subsequent paper (Cocker & Halsall, 1956) as (I), but the configuration at C(13) remained undetermined.



(I) R = CO₂H

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